

# **Vertical-Velocity Biases Caused by Reflectivity Fluxes**

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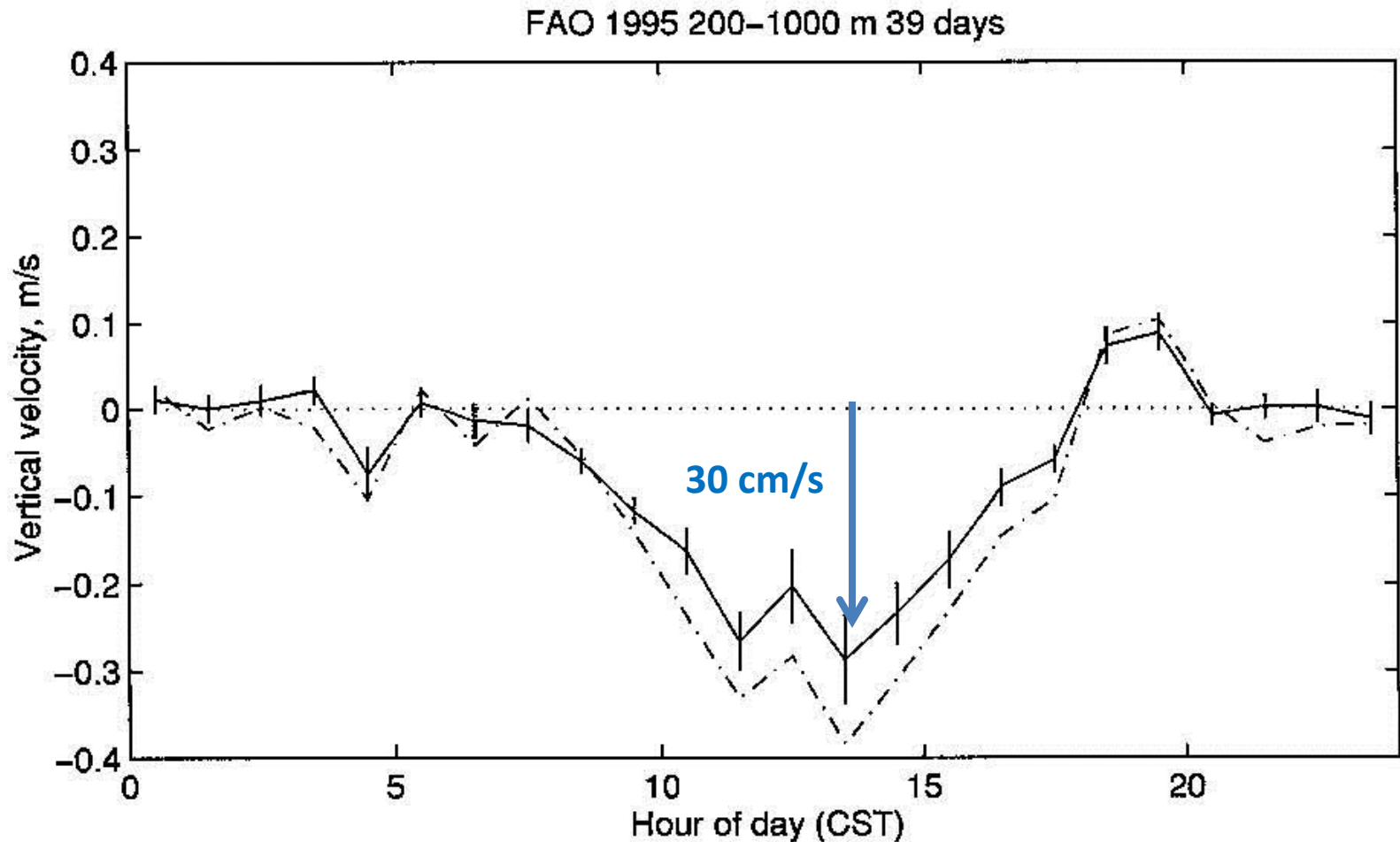
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# Overview

- Observations of vertical-velocity biases
- Intermittency fluxes: theory
- Intermittency fluxes: LES results
- Summary and a conclusion

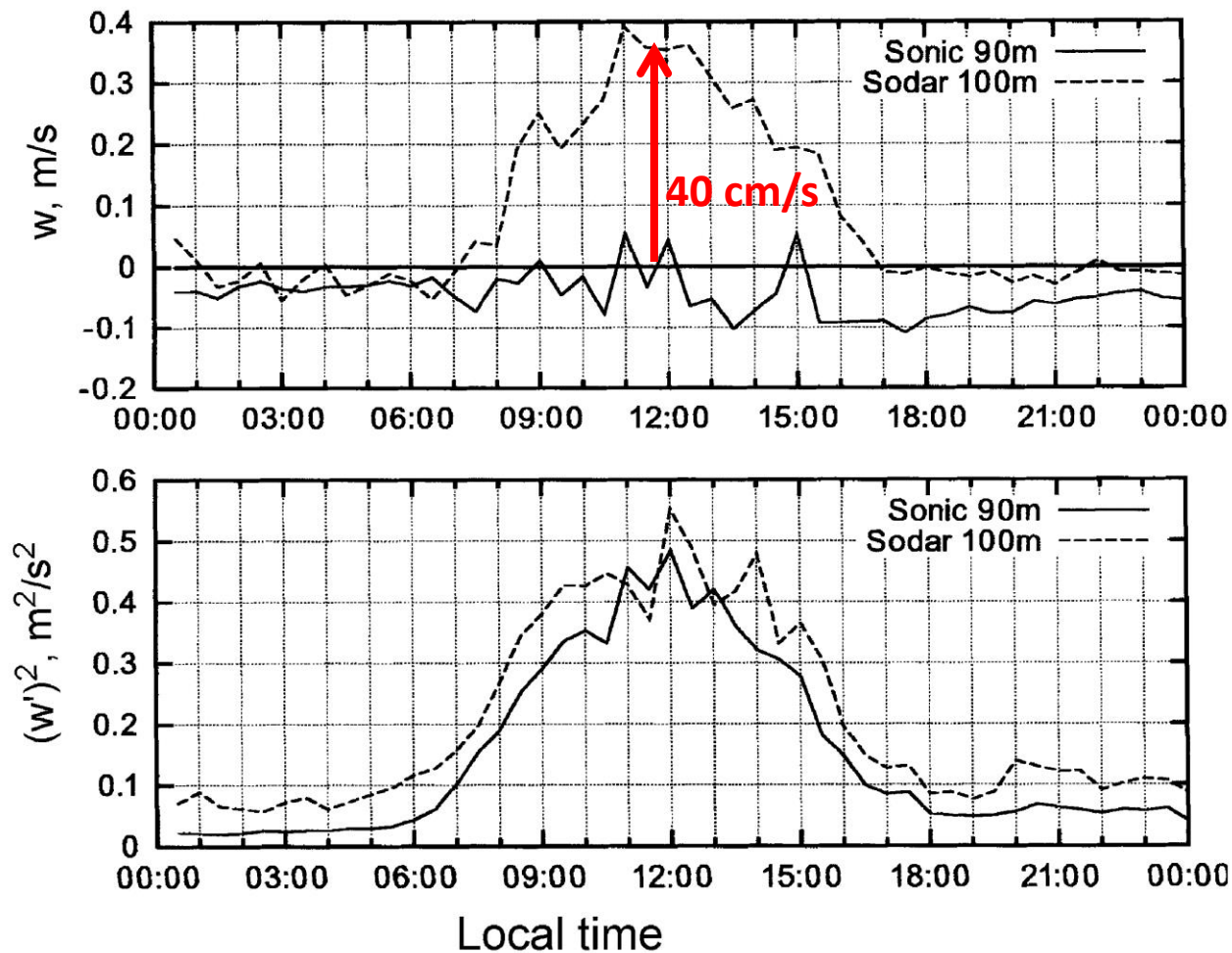
Observations:

Clear air radar wind profilers (**RWPs**) see mean **downward** wind (Angevine, 1997)



Observations:

**Sodars** see mean **upward** wind  
(Coulter and Kallistratova, 2004)



# Theory: intermittency biases of Doppler velocities

The vertical velocity observed with a vertically pointing radar windprofiler is

$$w_m = \frac{\langle \eta w \rangle}{\langle \eta \rangle}, \quad (1)$$

where  $w$  is the (true) vertical wind velocity,  $\eta$  is the clear-air radar reflectivity, and  $\langle \cdot \rangle$  is a space-time average over the radar's resolution volume and the radar's dwell time.

Now, allow  $\eta$  and  $w$  to be random variables, such that we can write them as sums of mean values and fluctuations:

$$\eta = \langle \eta \rangle + \eta' \quad (2)$$

and

$$w = \langle w \rangle + w'. \quad (3)$$

Then,

$$\langle \eta w \rangle = \langle (\langle \eta \rangle + \eta')(\langle w \rangle + w') \rangle = \langle \eta \rangle \langle w \rangle + \langle \eta' w' \rangle. \quad (4)$$

Inserting into (1) gives

$$w_m = \langle w \rangle + \frac{\langle \eta' w' \rangle}{\langle \eta \rangle}. \quad (5)$$

**Case 1:  $\eta$  and  $w$  are constant**

Then  $\eta'=0$  and  $w'=0$ , and

$$w_m = \langle w \rangle \quad (6)$$

(no bias).

**Case 2:  $\eta$  and  $w$  are random variables but uncorrelated**

Then  $\langle \eta' w' \rangle$ , and

$$w_m = \langle w \rangle \quad (7)$$

(again, no bias).

**Case 3:  $\eta$  and  $w$  are correlated random variables**

Then the reflectivity flux (or “intermittency flux”)  $\langle \eta' w' \rangle$  causes an **intermittency bias**

$$\Delta w = \frac{\langle \eta' w' \rangle}{\langle \eta \rangle}. \quad (8)$$

**Clear-air radar windprofiler:**

$$\eta = aC_n^2, \tag{9}$$

where  $a$  is a known constant and  $C_n^2$  is the (radio-wave) refractive-index structure parameter.

**Sodar:**

$$\eta = bC_T^2, \tag{10}$$

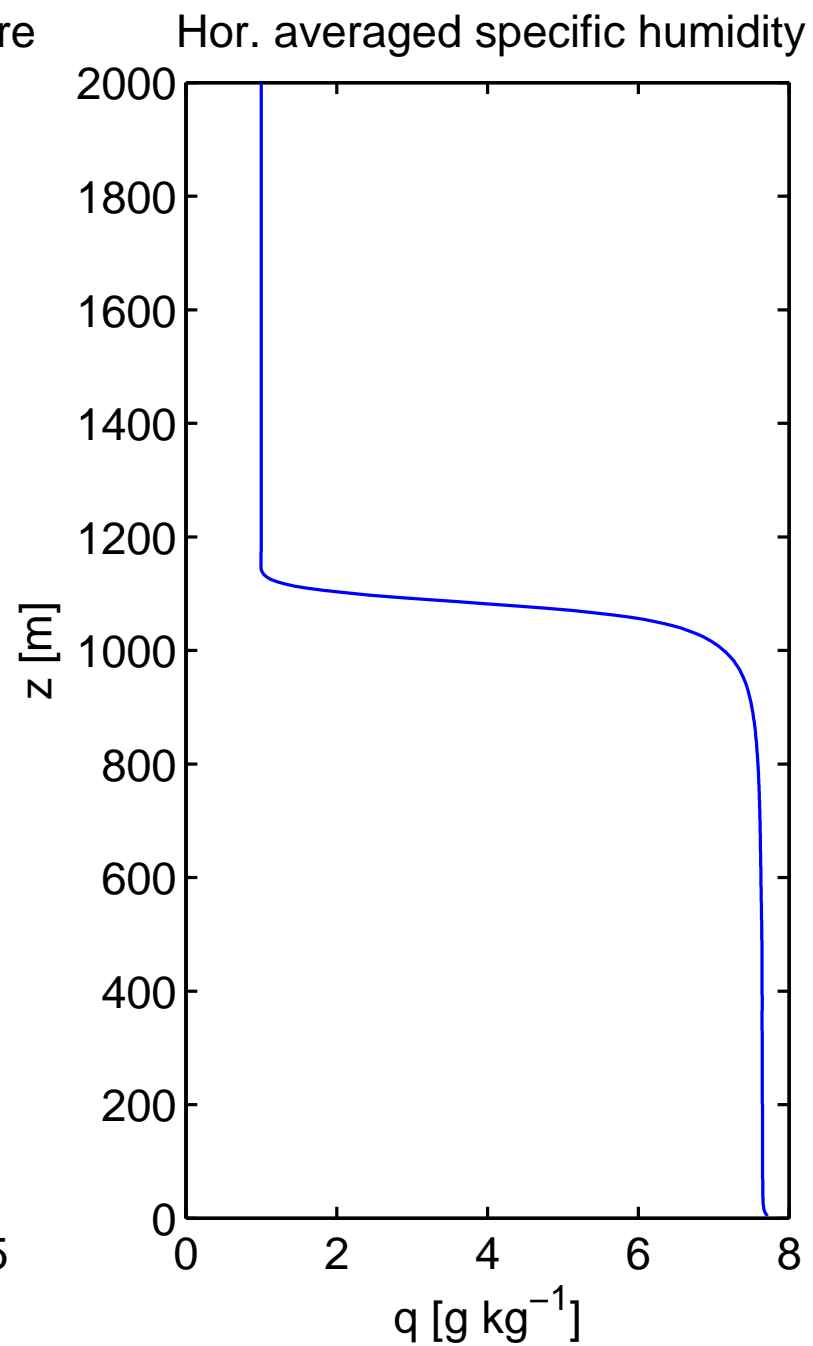
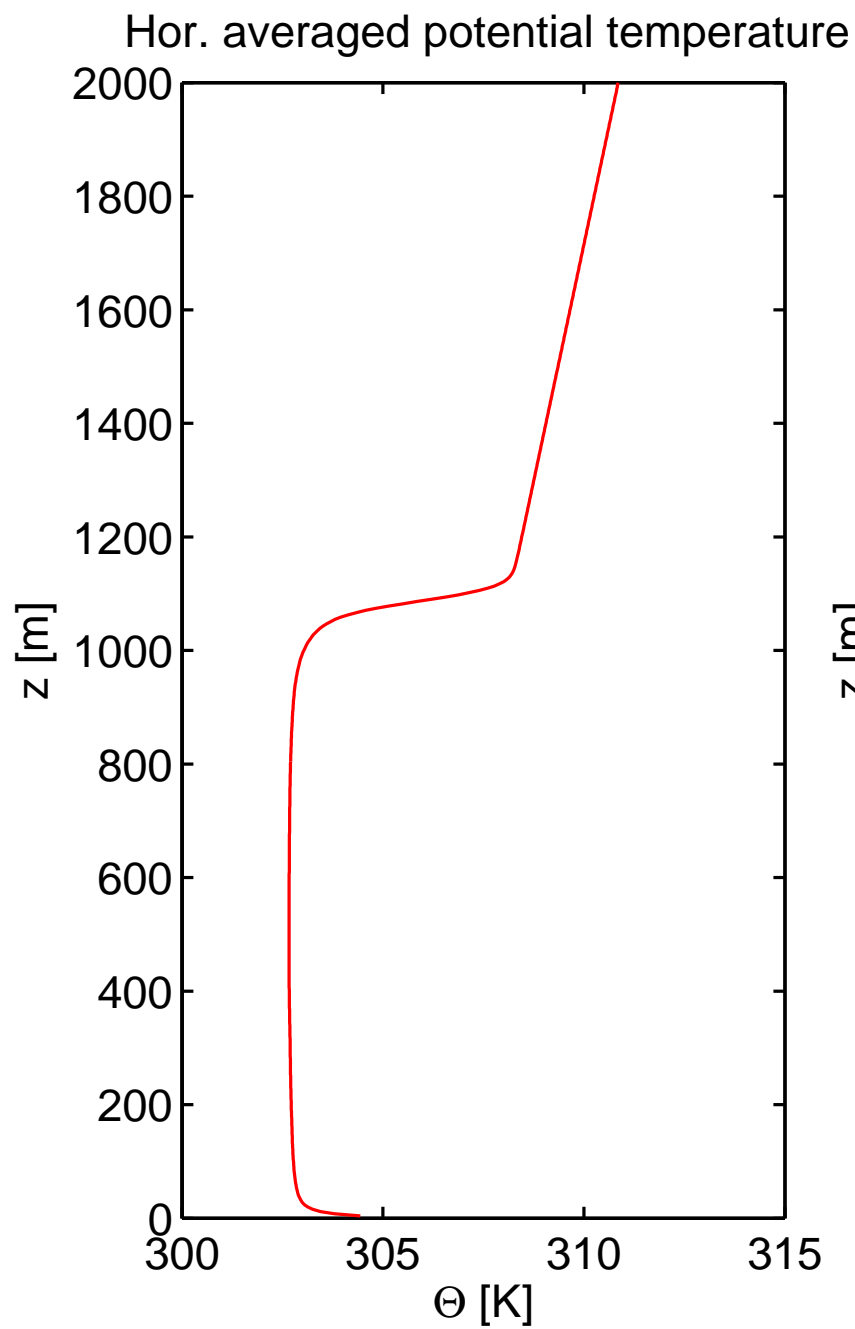
where  $b$  is also a known constant and  $C_T^2$  is the temperature structure parameter.

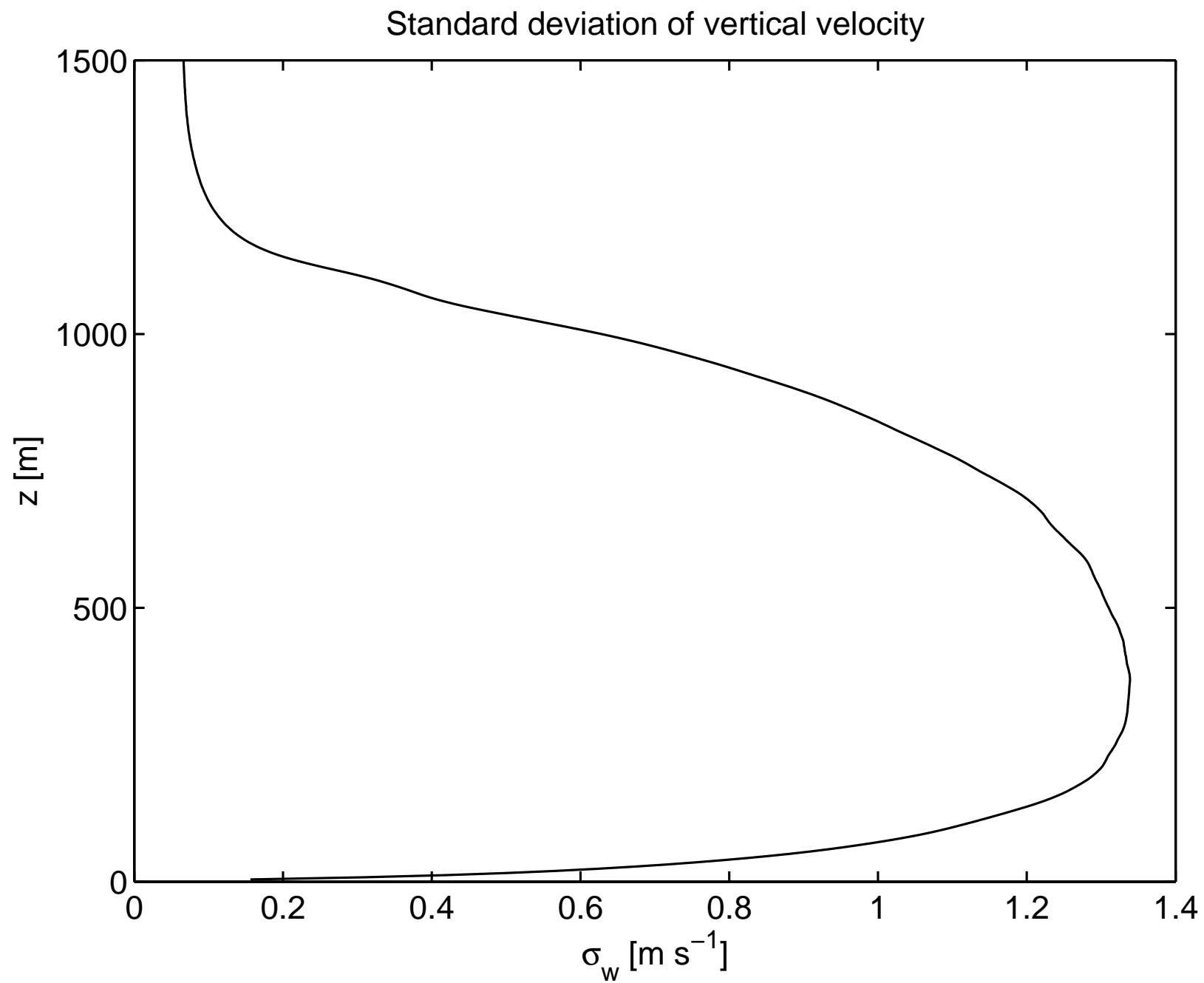
**Large-eddy simulation (LES)** can be used to evaluate mean values and fluctuations of  $C_n^2$  and  $C_T^2$  and their covariances with wind-velocity fluctuations.

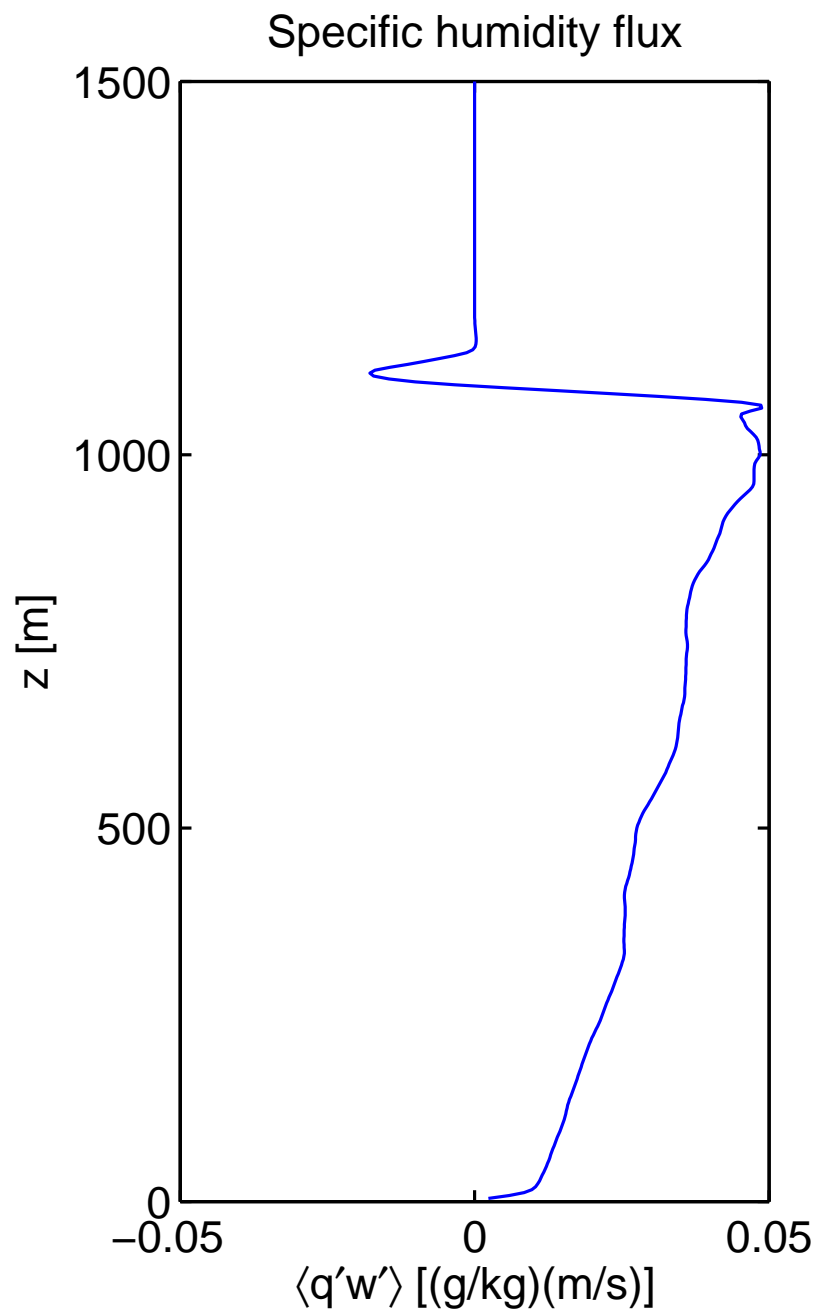
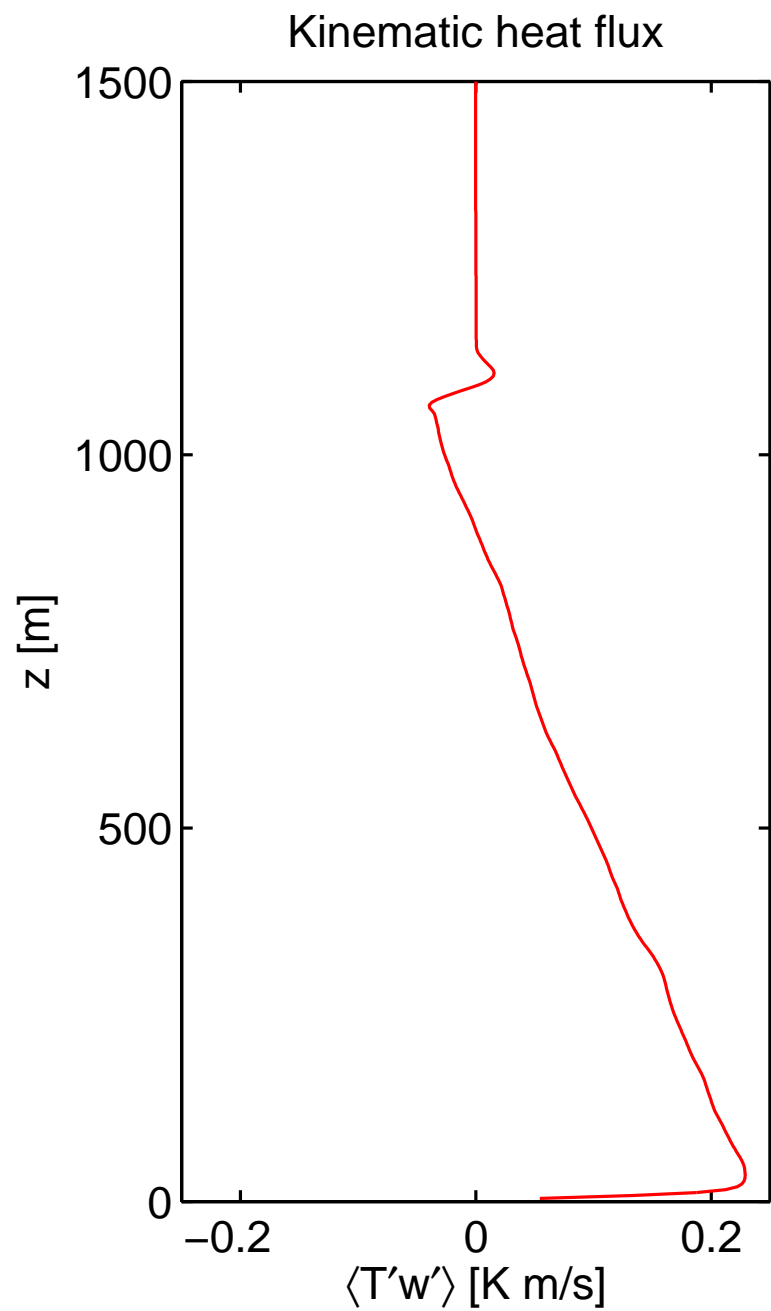
# **Large-eddy simulation (LES) of a convective boundary layer**

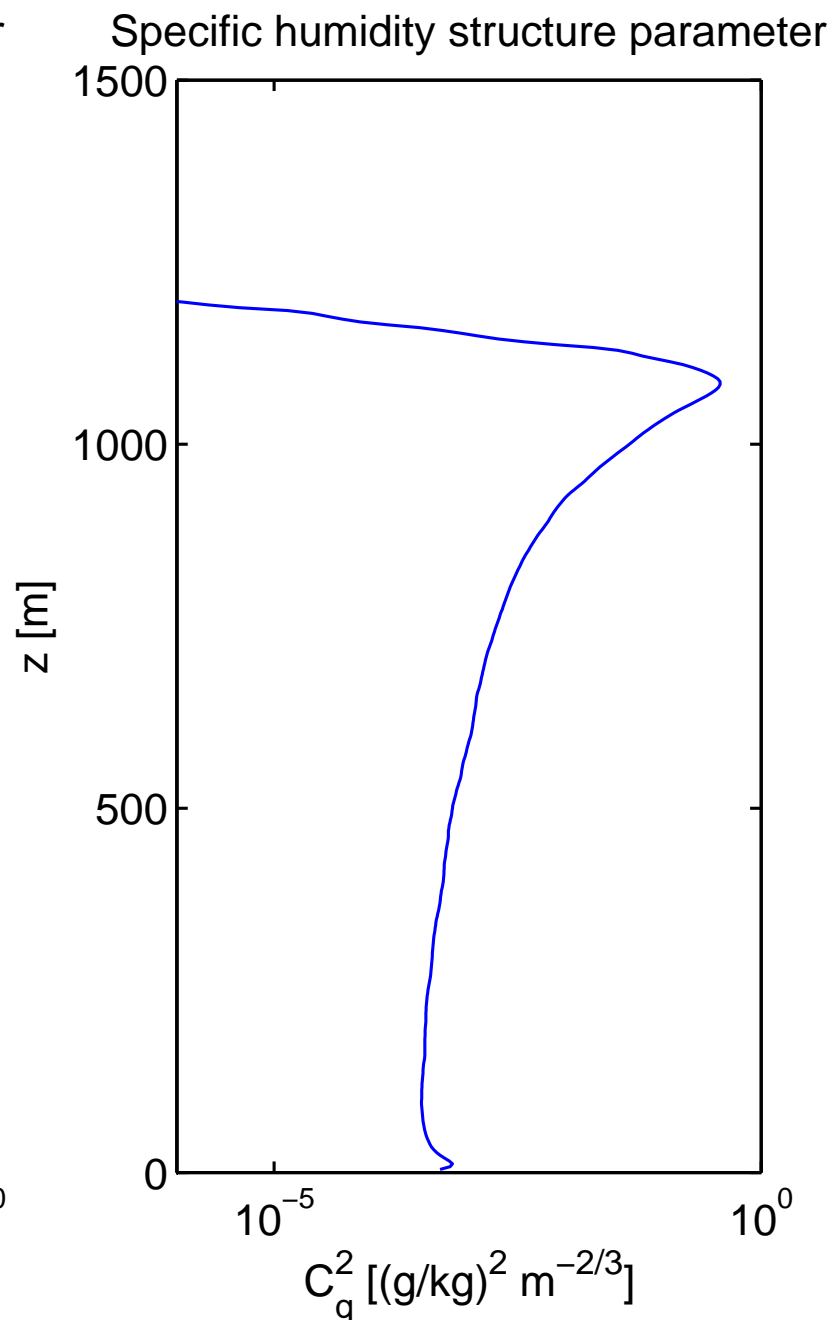
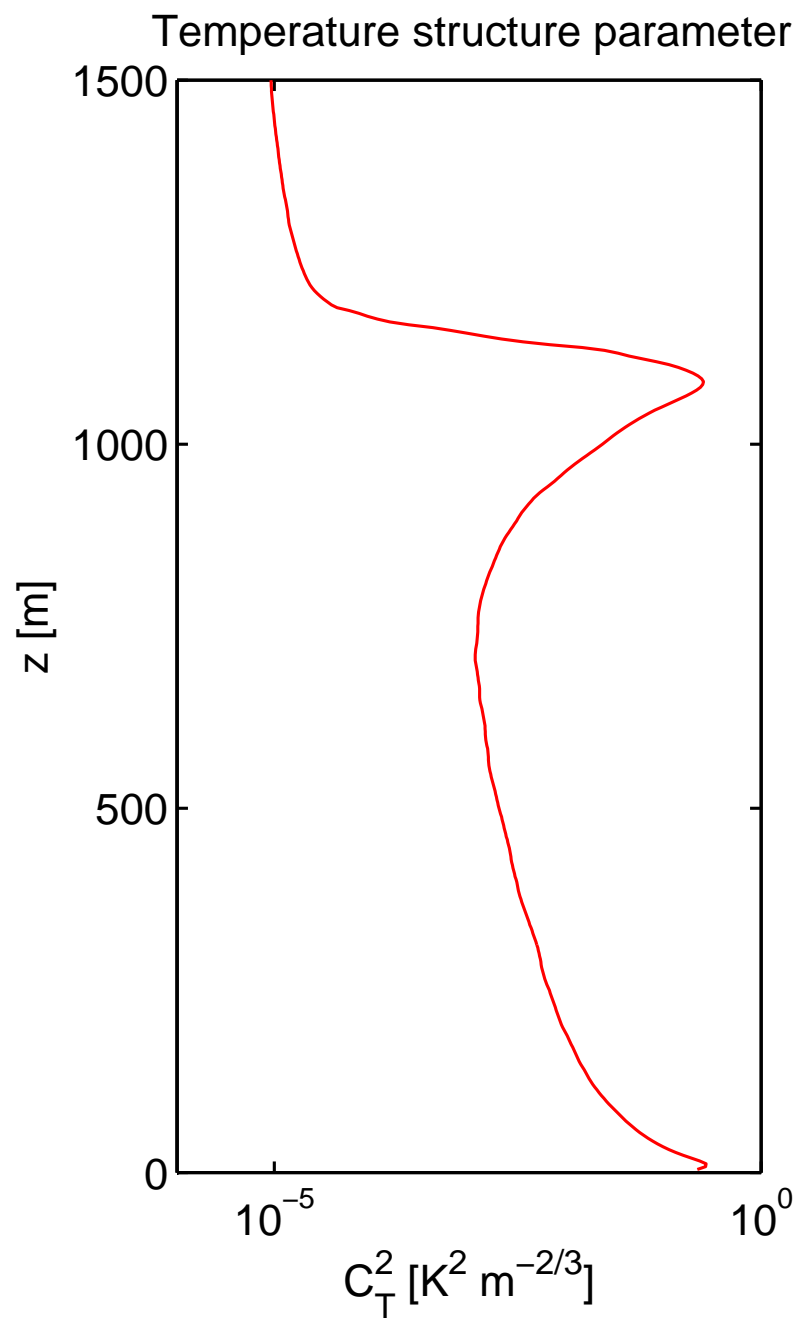
- Model domain: cube with 2000 m side length
- Grid size:  $512^3$
- Spatial resolution: 3.9 m in all three dimensions
- Mixed-layer height: 1100 m
- Kinematic heat flux:  $250 \text{ W m}^{-2}$

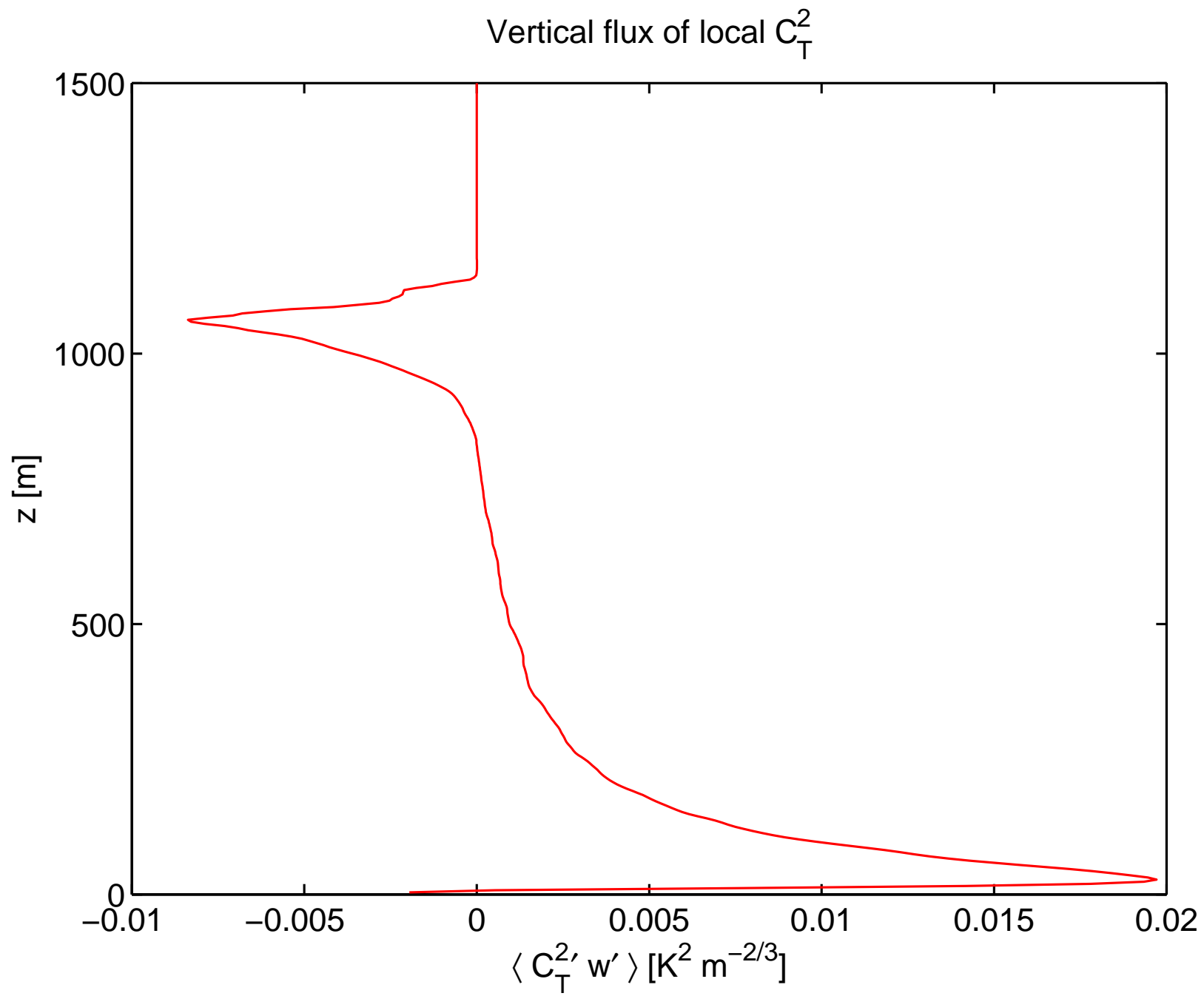


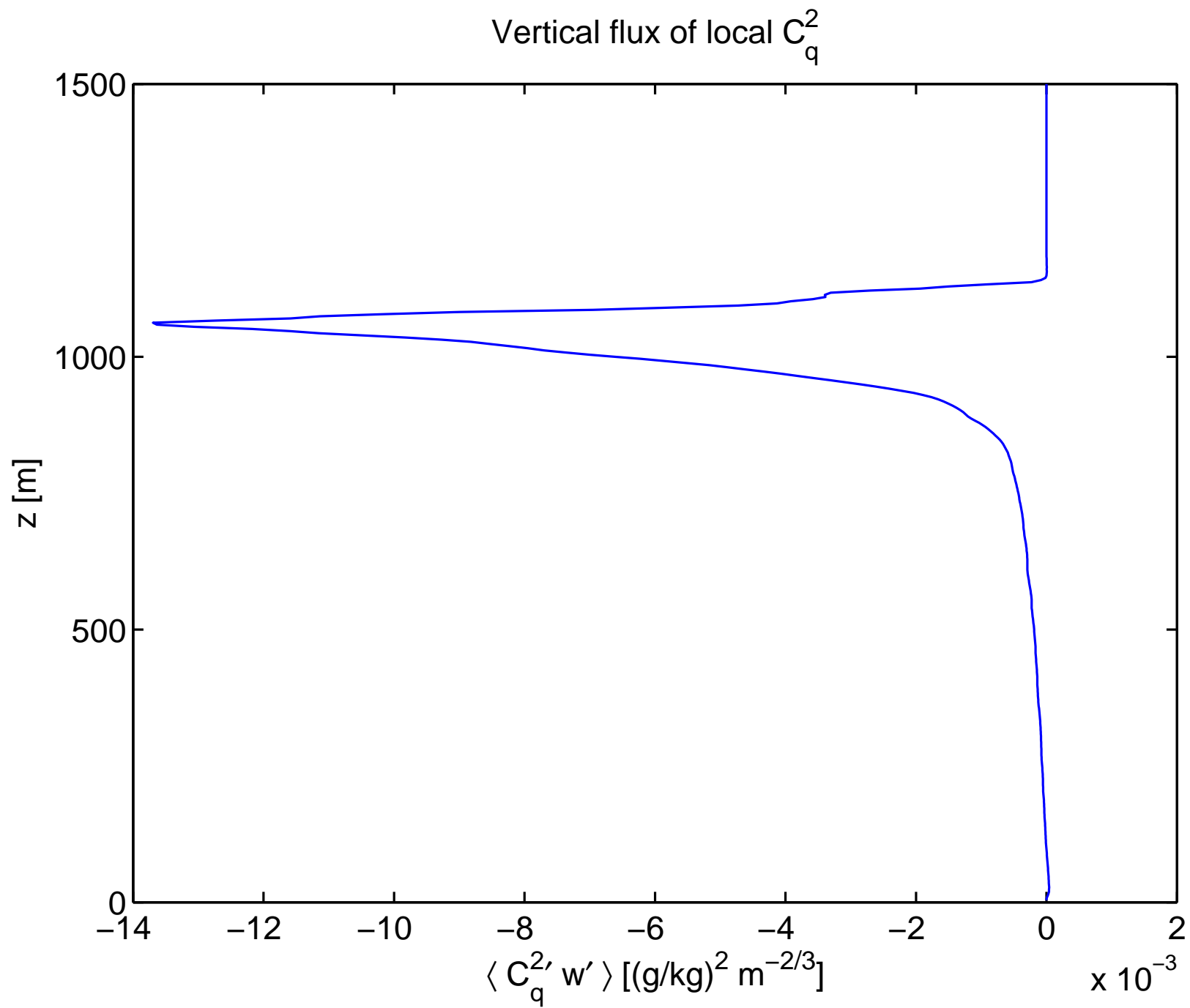




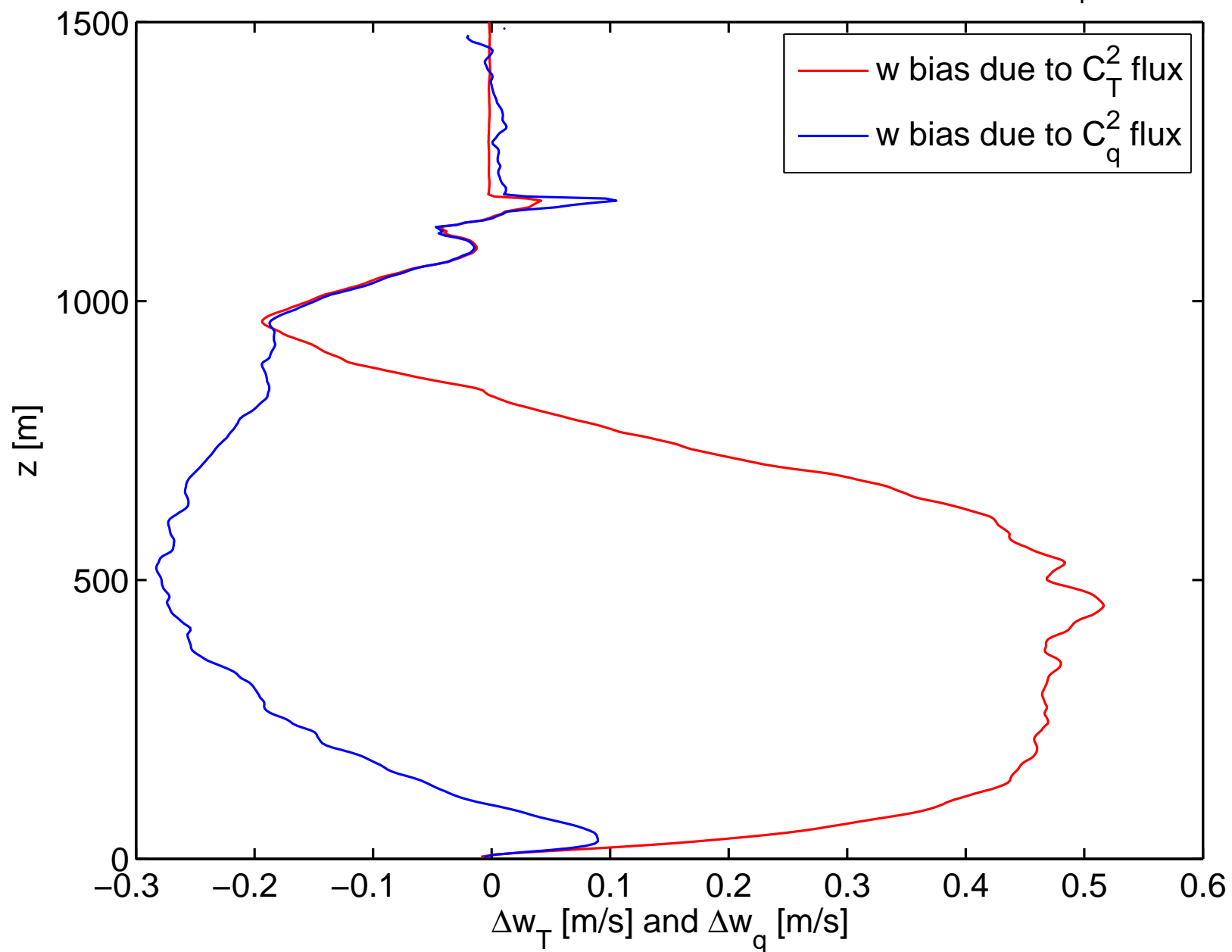








Doppler-velocity biases due to fluxes of local  $C_T^2$  and  $C_q^2$



# Summary and a conclusion

- RWP observations in the CBL show downward biases
- Sodar observations in the CBL show upward biases
- Theory explains how RWP/sodar velocity biases can be caused by “intermittency fluxes” of  $C_n^2$
- Intermittency fluxes retrieved from a large-eddy simulation of a canonical, “drying” CBL lead to “intermittency biases” that are qualitatively and quantitatively consistent with observations.
- It may be possible to use observations of Doppler-velocity biases in the CBL for the retrieval of vertical fluxes of  $C_T^2$  and  $C_q^2$ .